# **Atmospheric Chemistry**

Lecture 18

# **Excellent Tutorial Article on Tropospheric Gas-Phase Chemistry**

**TUTORIAL REVIEW** 

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## Gas-phase radical chemistry in the troposphere

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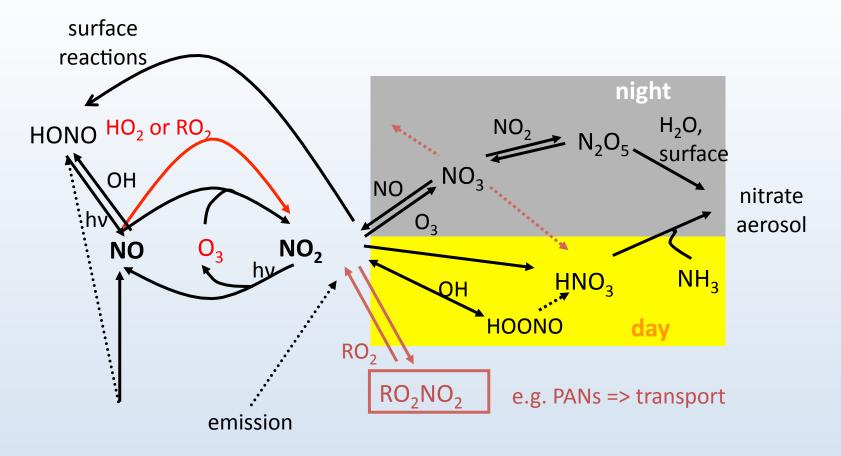
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Atmospheric free radicals are low concentration, relatively fast reacting species whose influence is felt throughout the atmosphere. Reactive radicals have a key role in maintaining a balanced atmospheric composition through their central function in controlling the *oxidative capacity* of the atmosphere. In this *tutorial review*, the chemistry of three main groups of atmospheric radicals  $HO_x$ ,  $NO_x$  and  $XO_x$  (X = Cl, Br, I) are examined in terms of their sources, interconversions and sinks. Key examples of the chemistry are given for each group of radicals in their atmospheric context.

## **Simplified NOx Chemistry in the Troposphere**



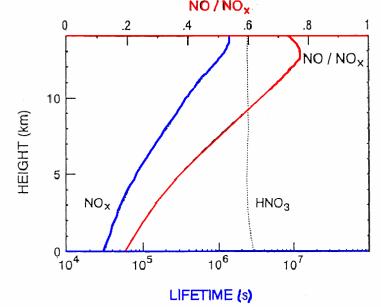
adapted from M. Jenkin

Nitrogen Oxides in the Troposphere, Andreas Richter, ERCA 2010

From an excellent review of the role of NOx in tropospheric chemistry by Andreas Richter obtained by googling "nox troposphere richter"

# Some facts on NO<sub>x</sub> in the Troposphere

- NO and NO<sub>2</sub> are rapidly converted into each other and are therefore combined to NO<sub>x</sub> = NO + NO<sub>2</sub>
- the ratio [NO] / [NO<sub>x</sub>] is about 0.2 at the surface but increases towards higher altitudes (temperature dependence of O<sub>3</sub> + NO reaction)
- the atmospheric lifetime of NO<sub>x</sub> is short close to the surface (hours) and increases towards higher altitudes (days)



- lifetime is longer in winter than in summer (lower [OH])
- the short lifetime results in little transport, both vertically and horizontally, at least in the form of NO<sub>x</sub>
   NO<sub>x</sub> is found close to its sources
- PAN has a long lifetime and can be transported and re-release NO<sub>x</sub> when temperature increases

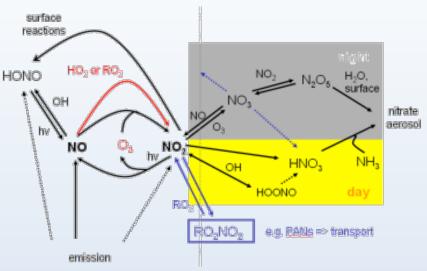
# The Role of NO<sub>x</sub> in Ozone Chemistry

#### **Background conditions**

- photolysis of NO<sub>2</sub> is only known way to produce O<sub>3</sub> in the troposphere
- O<sub>3</sub>, NO<sub>2</sub>, and NO are in photostationary state (*Leighton relationship*):

$$[O_3][NO_2] / [NO] = J_{NO2} / k_{O3 + NO}$$

 at very low [NO] / [O<sub>3</sub>] ratio, destruction of O<sub>3</sub> by HO<sub>2</sub> dominates over O<sub>3</sub> production



#### **Polluted conditions**

- if NO is oxidized to NO<sub>2</sub> by HO<sub>2</sub> or RO<sub>2</sub> instead of O<sub>3</sub>, ozone is catalytically formed by NO
- how much  $O_3$  can be formed in the presence of  $NO_x$  is eventually limited by the amount of CO,  $CH_4$ , and other hydrocarbons available
- at high  $NO_x$  concentrations,  $O_3$  levels are reduced by reaction with NO and  $NO_2$  in particular at night

## Net result for CO oxidation

#### **Can Produce Ozone**

$$CO + OH \rightarrow CO_2 + H$$

$$H + O_2 + M \rightarrow HO_2 + M$$

$$HO_2 + NO \rightarrow OH + NO_2$$

$$NO_2 + hv \rightarrow NO + O$$

$$O + O_2 + M \rightarrow O_3 + M$$

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$$CO + 2O_2 \rightarrow CO_2 + O_3$$

### **Can Consume Ozone**

$$CO + OH \rightarrow CO_2 + H$$

$$H + O_2 + M \rightarrow HO_2 + M$$

$$HO_2 + O_3 \rightarrow OH + 2O_2$$

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$$CO + O_3 \rightarrow CO_2 + O_2$$

**Depending on the NO<sub>x</sub> Concentration** 

# CO, NO, NO<sub>2</sub>, OH, HO<sub>2</sub> and J(O¹D) Measurements in Urban Summer Conditions

Note different scale for HO<sub>2</sub> measurements compared to OH measurements.

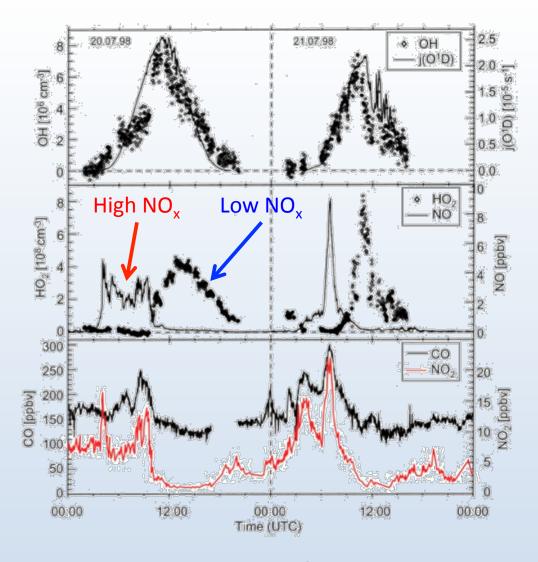


Fig. 8 Measurements of OH,  $HO_2$ ,  $j(O^1D)$ , NO,  $NO_2$  and CO during a summer urban campaign. The figure demonstrates the dramatic feedback of  $NO_x$  on  $HO_x$  radicals, the  $[HO_2]$  clearly suppressed at high  $NO_x$  while the [OH] remains largely unaffected.

from Monks, 2005

## **Net Ozone Production vs NO<sub>x</sub>**

- A: NO<sub>x</sub> concentrations very low; loss of ozone due to OH + O<sub>3</sub> and HO<sub>2</sub> + O<sub>3</sub> dominate
- B: NO<sub>x</sub> increasing leads to more efficient diversion of HC oxidation into ozone-producing reactions, e.g. converts NO to NO<sub>2</sub>
- C: Peak as production saturates; decrease in efficiency as conversion to organic nitrates removes some HC products before ozone production

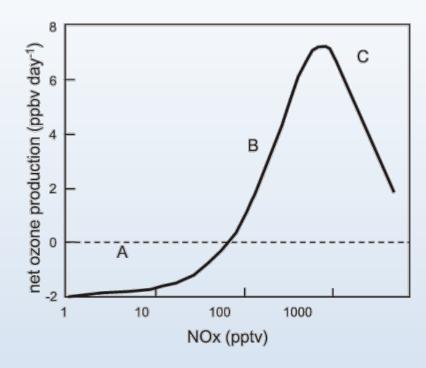


Fig. 9 Schematic representation of the dependence of the net ozone  $(N(O_3))$  production (or destruction) on the concentration of  $NO_x$ . The magnitudes reflect clean free tropospheric conditions.

## **Net Ozone Production vs NOx and VOCs**

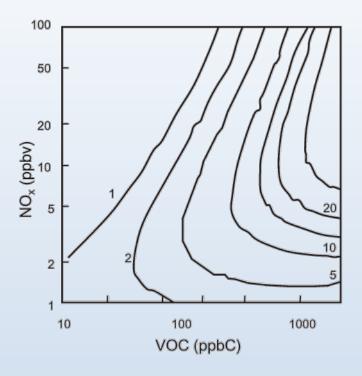


Fig. 10 Isopleths giving net rate of ozone production (ppb  $h^{-1}$ ) as a function of VOC (ppbC) and NO<sub>x</sub> (ppbv) for mean summer daytime meteorology and clear skies under urban conditions.<sup>16</sup>

- Note that NO<sub>x</sub> begins at 1 ppbv (1000 pptv), i.e. urban conditions
- Turnover point (as a function of NO<sub>x</sub>) increases as amount of VOCs increase

# Sources of NO<sub>x</sub> in the Troposphere

## Main sources of NO<sub>x</sub> (in Tg N / yr) are

•	fossil	fual	combustion
•	102211	Tue	COMBUSTION

fires

microbial soil emissions

lightning

oxidation of biogenic NH<sub>3</sub>

aircraft

stratosphere



6.7(3-10)

5.5(3.3-7.7)

2.0(1-4)

1.0(0.5-1.5)

0.5(0.5-0.6)

0.5(0.4-0.6)

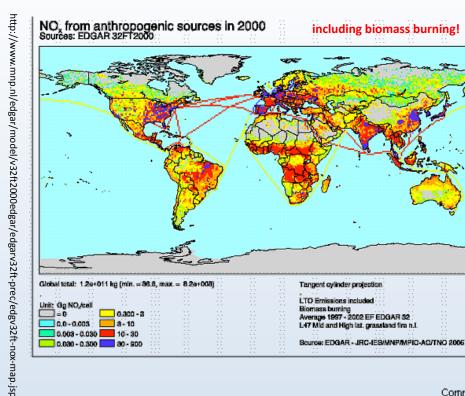






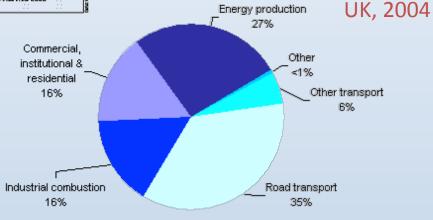






- anthropogenic emissions centered in a few industrialised areas
- largest emissions in cities and from power plants
- emissions per capita very unevenly distributes => future?

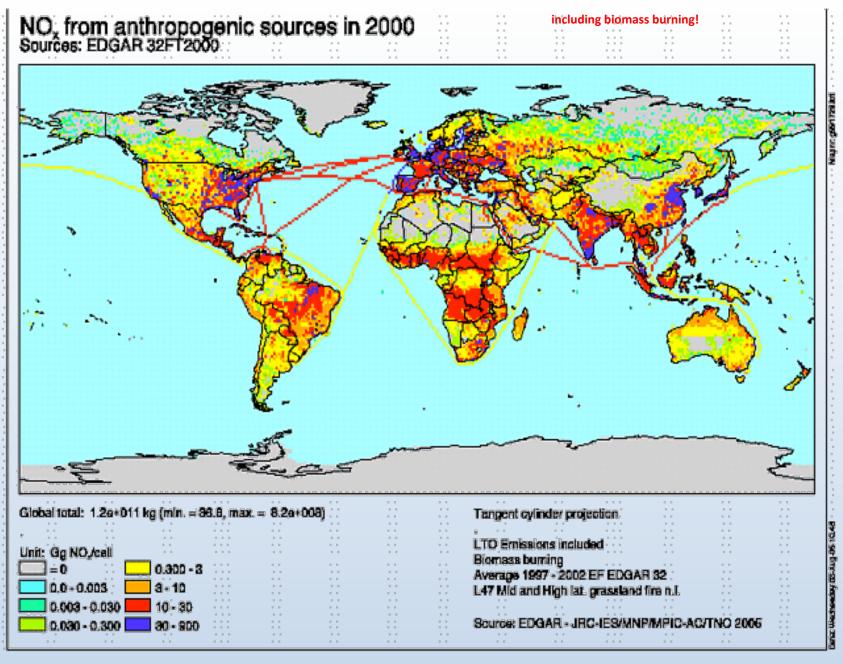
- road transport has large importance
- energy production is second, depending on energy mix



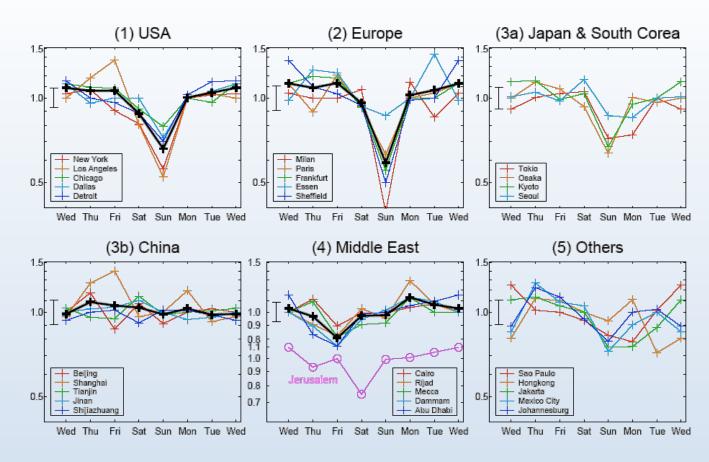
Total emissions of oxides of nitrogen = 1.67 million tonnes

Source: NAEI

-agency.gov.uk/commondata/103196/1162897?referrer=/yourenv/eff/1190084/air/1158715/1162725/



## Anthropogenic NO<sub>x</sub> Sources: Example



- Normalised tropospheric NO<sub>2</sub> columns retrieved from GOME satellite measurements show clear weekly cycle over industrialised areas
- anthropogenic NO<sub>x</sub> emissions dominate

# NO<sub>x</sub> from Lightning

at very high temperatures (> 2000 K)

$$O_2 + M \rightarrow O + O + M$$
  
 $O + N_2 \rightarrow NO + N$   
 $N + O_2 \rightarrow NO + O$   
(Zel'dovitch mechanism).

- lightning NO<sub>x</sub> is computed from the product of lightning dissipation energy and NO yield per Joule of discharge
- estimates have varied dramatically in the past:
   1.2 Tg ... 200 Tg N / yr





- lightning NO<sub>x</sub> in models often parameterised by cloud height or convective precipitation
- the relevance of lightning NO<sub>x</sub> is that it is injected in the upper troposphere, where ozone formation is very efficient

